

1 Do hurricanes cause significant interannual variability in the air-sea CO<sub>2</sub> flux of the subtropical  
2 North Atlantic?

3

4 Jennifer Koch, Galen A. McKinley\*, Val Bennington, and David Ullman

5 Atmospheric and Oceanic Sciences, University of Wisconsin-Madison, 1225 W. Dayton St,  
6 Madison, WI, 53706, Wisconsin, USA.

7

8 \*Corresponding author

9 Department of Atmospheric and Oceanic Sciences, University of Wisconsin-Madison, 1225 W.

10 Dayton Street, Madison, Wisconsin, 53706, USA

11 Tel +1 608 262 4817

12 Fax +1 608 262 0166

13 \* Email: gamckinley@wisc.edu

14

14  
15  
16  
17  
18  
19  
20  
21  
22  
23  
24  
25  
26  
27  
28

**1. Abstract**

Observations at Bermuda and in the Caribbean Sea indicate that hurricanes significantly influence surface ocean  $p\text{CO}_2$  ( $p\text{CO}_2^{\text{ocean}}$ ) and the air-sea  $\text{CO}_2$  flux at short time scales (Bates et al. 1998, Bates 2007, D’Asaro and McNeil 2007, Wanninkhof et al. 2007). We use a North Atlantic regional biogeochemical model to investigate causal mechanisms and to consider basin-wide impacts on interannual variability in air-sea  $\text{CO}_2$  fluxes. Consistent with previous observations, we find that enhanced wind speeds dominate the hurricane's effect on the flux, driving  $\text{CO}_2$  out of the ocean due to the negative air-sea gradient in  $p\text{CO}_2$  ( $p\text{CO}_2^{\text{atm}} < p\text{CO}_2^{\text{ocean}}$ ) that occurs in response to high sea surface temperatures (SSTs) during hurricane season (late summer / fall). Vertical mixing causes negative SST anomalies, which depress  $p\text{CO}_2^{\text{ocean}}$  but are not sufficient to reverse the air-sea gradient. Though hurricanes drive a local  $\text{CO}_2$  efflux, we find no evidence for a relationship between year-to-year variability in hurricane frequency and year-to-year variability in air-sea  $\text{CO}_2$  fluxes across the subtropical North Atlantic.

## 28 2. Introduction

29 The seasonal cycle of CO<sub>2</sub> exchange between the atmosphere and the subtropical North Atlantic  
30 Ocean is forced primarily by changes in the sea surface temperature (SST) due to seasonal heat  
31 fluxes (Takahashi et al. 2002, Bates 2007 (hereafter B07), Ullman et al. 2008). During the  
32 summer, pCO<sub>2</sub><sup>ocean</sup> rises above atmospheric pCO<sub>2</sub> and thus leads to an efflux of carbon dioxide to  
33 the atmosphere. The high wind speeds characteristic of hurricanes and tropical storms cause a  
34 rapid release of CO<sub>2</sub> from the ocean due to the negative summertime gradient in pCO<sub>2</sub> (i.e.  
35 pCO<sub>2</sub><sup>atm</sup> < pCO<sub>2</sub><sup>ocean</sup>). A deepening of the MLD also occurs with hurricane passage, causing a  
36 decrease in the SST of about 2-3°C (Bates et al. 1998 (hereafter B98), Perrie et al. 2004, D'Asaro  
37 and McNeil 2007, Wanninkhof et al. 2007). Despite the sometimes-large decrease in pCO<sub>2</sub><sup>ocean</sup>  
38 due to reduced SSTs, pCO<sub>2</sub><sup>ocean</sup> remains higher than pCO<sub>2</sub><sup>atm</sup>, so the wind speed enhancement  
39 causes a strong CO<sub>2</sub> efflux (B98, D'Asaro and McNeil 2007, Wanninkhof et al. 2007).

40

41 Though their local effects have been clearly illustrated, the role that hurricanes play in the annual  
42 air-sea exchange of CO<sub>2</sub> in the subtropics is not entirely clear. Following the work of B98 who  
43 estimated a net global flux up to 0.51 PgC/yr due to hurricanes, B07 estimates a global flux of  
44 0.04-0.08 PgC/yr due to hurricanes. However, these fluxes could simply be a short-term  
45 enhancement of the summer/fall outgassing. Would the flux have occurred anyway, albeit more  
46 gradually, without hurricanes? In this paper, we consider the impacts of hurricanes on surface  
47 ocean biogeochemistry near Bermuda and on the interannual variability in air-sea CO<sub>2</sub> fluxes in  
48 the subtropical North Atlantic from 25-40N using an ocean biogeochemical model.

49

49 **3. Model**

50 As described by Bennington et al. (2008) and Ullman et al. (2008), the MIT Ocean General  
51 Circulation Model (Marshall et al. 1997) is regionally configured for the North Atlantic between  
52 20°S and 81.5°N, with a horizontal resolution of 0.5° latitude and 0.5° longitude. The model has  
53 23 vertical levels with a resolution of 10m thickness at the surface increasing to 500m thickness  
54 for depths greater than 2200m. Relaxation to hydrographic and biogeochemical tracers occurs  
55 along regional boundaries. The model is forced with daily fields from the NCEP/NCAR  
56 Reanalysis I (Kalnay et al. 1996) for 1992-2006 and SSTs are relaxed (2 week timescale) to  
57 1992-2006 satellite-based estimates (Reynolds et al. 2007). After an 81-year physical spin up, the  
58 biogeochemical model was initialized with climatological DIC, alkalinity and nutrients, low  
59 values of phytoplankton and zooplankton, and atmospheric pCO<sub>2</sub> fixed at 356 ppm (1992 Mauna  
60 Loa mean), and was run for 70 years. For the main run, atmospheric pCO<sub>2</sub> data is interpolated  
61 from monthly Mauna Loa observations (Keeling et al. 2001) and is adjusted for local surface  
62 atmospheric pressure. The biogeochemical model includes carbonate chemistry and nutrient  
63 cycling using an ecosystem model of intermediate complexity (Dutkiewicz et al. 2005) that  
64 models a pelagic ecosystem with one zooplankton class and two phytoplankton classes: diatoms  
65 and “small” phytoplankton. The flux of carbon dioxide between the ocean and atmosphere is  
66 parameterized following Wanninkhof (1992), shown by McNeil and D’Asaro (2007) to be  
67 appropriate when considering hurricanes. We define a positive flux to be out of the ocean.

68

69

69        4. Results

70    4a. Local CO<sub>2</sub> flux response to 1995 hurricanes at Bermuda

71    For comparison to B98 we begin by asking: How well does the NCEP/NCAR Reanalysis capture  
72    the hurricanes of 1995? Of the three storms, Hurricane Felix passed closest to Bermuda, coming  
73    nearest on August 14-15. Figure 1a demonstrates a peak in the wind speed on these days,  
74    indicating that the presence of Felix was replicated in the Reanalysis. There was a low-pressure  
75    center (not shown), with the lowest pressure on August 15<sup>th</sup> 1995 at 988.4 mb. The presence of  
76    Luis (September 8-9) and Marilyn (September 18-19) are also notable in the wind speeds.

77

78    In comparison to the data, modeled MLDs are reasonable across both the seasonal cycle and  
79    during the hurricane season. The model responds to hurricane passage with increases in mixed  
80    layer depths (MLD) of 5-20m (Figure 1b). Increased mixing causes SSTs (Figure 1c) to drop  
81    during the passage of each hurricane, rebounding slowly in the following weeks (B98). pCO<sub>2</sub><sup>ocean</sup>  
82    decreases with the passage of the hurricanes (Figure 1d). In order to better understand this drop  
83    in pCO<sub>2</sub><sup>ocean</sup>, it is split into the components driven by temperature changes (pCO<sub>2</sub>-T) and by non-  
84    temperature effects (pCO<sub>2</sub>-nonT) (Takahashi et al. 2002, see Supplementary text). With this split,  
85    is clear that SST depression drives the drop in pCO<sub>2</sub><sup>ocean</sup> (Figure 1d). With Felix, there is also a  
86    small decrease in pCO<sub>2</sub><sup>ocean</sup> due to non-temperature causes. Analysis of the surface DIC budget  
87    (Section 4b and Supplementary text) indicates this drop is due to dilution by large freshwater  
88    inputs. Despite the SST-driven drop, pCO<sub>2</sub><sup>ocean</sup> is still significantly above the pCO<sub>2</sub> of the  
89    atmosphere (pCO<sub>2</sub><sup>atm</sup>). Thus, the wind speed enhancement drives large CO<sub>2</sub> efflux anomalies  
90    (Figure 1e). In the days or weeks following the hurricane passage, there is a depression in the  
91    CO<sub>2</sub> flux, coincident with the lowered SST and pCO<sub>2</sub><sup>ocean</sup> (Wanninkhof et al., 2007).

92

93 The magnitudes of modeled CO<sub>2</sub> fluxes during the 1995 hurricanes are similar to those described  
94 by B98 for Luis and Marilyn, but substantially less for Felix. This is due in part to the model  
95 being forced with daily winds, which do not fully capture a hurricane's intensity. For example,  
96 Felix had maximum sustained winds of 40-45 m/s (B98), with maximum wind speeds at  
97 Bermuda of 26.2 m/s (B07), while the daily maximum from NCEP/NCAR Reanalysis at  
98 Bermuda is only 14.5 m/s (Figure 1a). NCEP maximum wind speeds for Luis (14.1 m/s) and  
99 Marilyn (13.5 m/s) were closer to those observed at Bermuda (19.1 m/s, 19.5 m/s, respectively).  
100 Modeled CO<sub>2</sub> fluxes during the two days of Felix, Luis, and Marilyn were 32, 18, and 15 (total  
101 65) mmol/m<sup>2</sup>, respectively, out of a total summer / fall efflux of 349 mmol/m<sup>2</sup> (a sum of all days,  
102 with flux from ocean to atmosphere in Figure 1d). B98 used observations to calculate fluxes of  
103 108, 21, and 18 (total 145) mmol/m<sup>2</sup>, and estimate a total summertime efflux of 331 mmol/m<sup>2</sup>.  
104 Because the model underestimates the hurricane fluxes and overestimates surface ocean pCO<sub>2</sub> in  
105 summer and fall (Bennington et al. 2008), and thus overestimates the total efflux in this period,  
106 the percentage of the total summertime efflux of carbon dioxide caused by the three hurricanes  
107 was 19% for the model, compared to 44% estimated by B98.

108

109 Though the daily winds forcing the model do not grossly underestimate hurricane winds at  
110 Bermuda, these winds do cause model-estimated hurricane-driven CO<sub>2</sub> fluxes to be substantially  
111 too small. The quadratic dependence of gas exchange on wind speed following the Wanninkhof  
112 (1992) parameterization we use is a primary reason (Wanninkhof et al. 2007). Averaging of  
113 observations in both time and space in the creation of the daily NCEP Reanalysis product  
114 contributes to these muted wind speeds. A model with greater horizontal and temporal resolution

115 would certainly improve our representation of the surface ocean response to hurricanes, but  
116 forcing data with appropriate resolution is not available for long enough to allow simulation of  
117 the interannual changes of central interest to this study – for example, QuickScat satellite winds  
118 are only available from 1999. Thus, for the purpose of this study, we illustrate that the model is  
119 able to reasonably capture the surface ocean response to all three 1995 hurricanes near Bermuda,  
120 and thus is a reasonable tool for study of the local biogeochemical response to hurricanes  
121 (section 4b) and variability in basin-averaged CO<sub>2</sub> fluxes (section 4c,d).

122

#### 123 4b. Hurricane impacts on the surface ocean biogeochemistry

124 Modeled phosphate (PO<sub>4</sub>) and DIC profiles, and MLDs for one day prior to and one day post  
125 hurricane passage are presented in Figure 2, along with monthly data profiles from BATS. The  
126 model nutricline sits slightly too high in the water column, surface DIC is too high, and the DIC  
127 vertical gradient is too weak, all consistent with low productivity and high summer pCO<sub>2</sub><sup>ocean</sup> as  
128 discussed above (see also Supplementary text), and may be due to a lack of subsurface to surface  
129 nutrient transfer in unresolved mesoscale eddies (Ledwell et al., 2008). These problems with the  
130 model mean state do likely degrade the representation of the biogeochemical response to  
131 hurricanes.

132

133 The modeled surface DIC response to the hurricane is not consistent from storm to storm. There  
134 is a large decrease in surface DIC with Felix, and moderate increases during Luis and Marilyn.  
135 Detailed model analysis (see Supplementary material) indicates that DIC loss during Felix is due  
136 to freshwater inputs and thus dilution of DIC. During Luis and Marilyn, the net DIC gain is due  
137 to vertical and horizontal advection and mixing, as well as net evaporation. For none of the three

138 hurricanes is a net loss of DIC due to the air-sea flux a dominant term in the surface ocean DIC  
139 budget. Changes in the deeper (>50m) profiles of both PO<sub>4</sub> and DIC suggest upward heaving of  
140 the seasonal thermocline in response to the hurricane.

141

142 Consistent with B98, storm-induced mixing in the model does not penetrate into the nutricline,  
143 bring nutrients to the surface, and stimulate a biological response. However, there have been  
144 satellite observations of enhanced surface ocean chlorophyll in the wake of a hurricane,  
145 something this model does not capture. Mixing of the subsurface chlorophyll maximum to the  
146 surface is a likely reason for this feature. The model's inability to capture this is due, at least in  
147 part, to the model's low productivity in the subtropical gyre and associated weak subsurface  
148 chlorophyll maximum. Additionally, model horizontal resolution is too coarse to capture the  
149 mesoscale and submesoscale processes that can drive strong local upwelling and mixing.

150

#### 151 4c. Basin-wide impact on carbon dioxide flux

152 In this section, we use our North Atlantic model to test the basin-scale extrapolations by B98 and  
153 B07. They conclude that the year-to-year differences in hurricane frequency and intensity might  
154 be an important mechanism for controlling interannual variability in air-sea CO<sub>2</sub> exchange  
155 globally between 40S and 40N. Neither our study or theirs considers impacts north of 40N,  
156 where hurricanes may drive enhanced carbon uptake due to the reversed air-sea pCO<sub>2</sub> gradient in  
157 summer. Furthermore, we focus on the subtropical region, 25-40N.

158

159 Data from National Hurricane Center (NHC) was compiled to locate hurricanes and tropical  
160 storms that passed north of 25N in the North Atlantic from 1992 to 2006. The number of

161 hurricanes per year is displayed in Figure 3a. The apparent increasing trend is not well defined  
162 by a linear trend, with an  $r^2$  value of 0.36 ( $p < 0.05$ ). However, the net North Atlantic  $\text{CO}_2$  sink for  
163 25-40N is increasing with a steady, significant trend ( $r^2 = 0.65$ ,  $p < 0.01$ , Figure 3b), due primarily  
164 to the anthropogenically-driven increase in  $\text{pCO}_2^{\text{atm}}$  (Ullman et al. 2008). Figure 3c compares the  
165 annual sum of the local  $\text{CO}_2$  flux occurring along the path of the hurricane or tropical storm  
166 (“along-track”) to the number of hurricanes. The trend is reasonably well defined by a linear  
167 relationship ( $r^2 = 0.56$ ,  $p < 0.1$ ), as expected from the above analysis of three hurricanes from 1995  
168 at Bermuda. Locally,  $\text{CO}_2$  efflux increases with the number of hurricanes. The fact that we  
169 consider only storm number suggests that it is not necessary to describe variation in storm size  
170 and strength to capture the first-order effects on local  $\text{CO}_2$  fluxes.

171  
172 If we now consider the  $\text{CO}_2$  flux across the basin between 25-40N, we find there is not a  
173 statistically significant relationship between the number of hurricanes and the annual carbon flux  
174 to the ocean ( $r^2 = 0.23$ ,  $p < 0.1$ , Figure 3d). In other words, the model suggests that variation in the  
175 number of North Atlantic hurricanes and tropical storms does not have a clear relationship to the  
176 magnitude of the  $\text{CO}_2$  sink in the subtropical North Atlantic. This finding is contrary to the  
177 suggestion of B98 and B07. It suggests that the large local fluxes occurring during a hurricane do  
178 not drive a significant additional efflux of  $\text{CO}_2$  from the ocean, but instead only temporarily  
179 enhance the process of summer / fall outgassing that is already underway.

180

181 4d. Are model errors likely to impact our conclusions?

182 This model underestimates the magnitude of the local surface ocean carbon cycle response to  
183 hurricanes, largely because of the use of daily-mean winds and the one-half degree resolution.

184 The model also overestimates the summertime air-sea gradient in pCO<sub>2</sub>. These two effects  
185 partially counteract each other in the calculation of the flux, but the result remains an  
186 underestimate (section 4a). Could these model flaws modify our conclusion that there is a  
187 negligible impact of hurricanes on interannual variability in the subtropical CO<sub>2</sub> sink? To test  
188 this, we consider the modeled maximum annual along-track flux for all hurricanes (223 mmol  
189 C/m<sup>2</sup>, Figure 3c) and assume it is an underestimate by more than a factor of 2; thus we will use  
190 500 mmol C/m<sup>2</sup> in our calculation. Note that this is more than 3 times the flux due to all 1995  
191 hurricanes (145 mmol C/m<sup>2</sup>) estimated by B98. If we assume this annual flux applies over all of  
192 a canonical 250km diameter hurricane, the largest size used by B98, the total flux is  $2.95 \times 10^{-4}$   
193 PgC/yr, only 0.26-0.44% of the annual net carbon flux estimate for the years 1992-2006 (-0.066  
194 to -0.114 PgC/yr, Figure 3b) for the subtropical North Atlantic from 25-40N. If we add to this  
195 tropical force winds extending to 750km and causing a flux of 150 mmolC/m<sup>2</sup>, the total flux is  
196  $1.00 \times 10^{-3}$  PgC/yr, or 0.87-1.5% of the total annual sink. We conclude that even if the model  
197 was able to more accurately simulate the local impact of hurricanes, we would still not find them  
198 to have a significant impact on year-to-year variability in the subtropical North Atlantic carbon  
199 sink.

200

201 Our findings are in contrast to previous works, particularly B98, which we attribute in part to the  
202 methodology used in their basin-scale extrapolation. They use very large hurricane-driven fluxes,  
203 values that are generally much larger (by 100 times or more) than those estimated from  
204 observations at Bermuda (see Supplementary text). Our modeled maximum annual hurricane-  
205 driven flux across the North Atlantic (25-40N, 223 mmol C/m<sup>2</sup>) is only 3.4 times the maximum  
206 modeled annual flux at Bermuda (65 mmol C/m<sup>2</sup>, 1995, section 4a).

207

207 5. Summary  
208

209 We have evaluated the impact of hurricanes on the carbon cycle of the subtropical North Atlantic  
210 using an ocean biogeochemical model that successfully captures observed seasonal patterns of  
211 MLD, SST and pCO<sub>2</sub> variations at Bermuda and also captures, to a reasonable degree, the effects  
212 of hurricane passage on these quantities. Reduced pCO<sub>2</sub> in the wake of three 1995 hurricanes is  
213 mostly due to the drop in SST that occurs with increased mixing, which supplies cool subsurface  
214 waters to the surface. Though the high winds associated with the storms cause an intense,  
215 temporary efflux of CO<sub>2</sub>, this drop in SST and corresponding reduced pCO<sub>2</sub> in the wake of the  
216 storms likely compensate for the enhanced flux during the storms. Nutrient supply to the surface  
217 with vertical mixing is weak because the mixing enhancement by the storm does not reach into  
218 the nutricline. Changes in DIC have little impact on surface ocean pCO<sub>2</sub>, and the changes that do  
219 occur are driven physically, not biologically. The model suggests that, at the scale of the whole  
220 subtropical North Atlantic, 25-40N, the frequency of hurricanes and tropical storms is not  
221 associated with significant changes in the net annual carbon sink from 1992-2006.  
222

222 Acknowledgements

223

224 We thank Stephanie Dutkiewicz for fruitful collaboration on model development and analysis;  
225 and Nick Bates for data and commentary; Rik Wanninkhof and an anonymous reviewer for their  
226 insightful comments; and the many scientists contributing to the BATS timeseries. We thank  
227 NASA for funding (CARBON/04-0300-0228).

228 References

229

230 Bates, N. R. A. H. Knap, and A. F. Michaels, (1998), Contribution of hurricanes to local and

231 global estimates of air-sea exchange of CO<sub>2</sub>. *Nature* 395.6697: 58-61.

232 Bates, N. R. (2007) Interannual variability of the oceanic CO<sub>2</sub> sink in the subtropical gyre of the

233 North Atlantic Ocean over the last 2 decades. *J. Geophys. Res.*, 112, C09013, doi:

234 10.1029/2006GC003759.

235 Bennington, V, G.A. McKinley, S. Dutkiewicz, D. Ullman (2008) What does chlorophyll

236 variability tell us about export and CO<sub>2</sub> flux variability?, *Global Biogeochem. Cycles, in*

237 *revision*.

238 D'Asaro, E. D., and C. McNeil (2007), Air-sea gas exchange at extreme winds speeds measured

239 by autonomous oceanographic floats, *J. Marine systems*, 66, 92-109.

240 Dutkiewicz, S., M. J. Follows, and P. Parekh (2005), Interactions of the iron and phosphorus

241 cycles: A three-dimensional model study, *Global Biogeochem. Cycles*, 19, GB1021,

242 doi:10.1029/2004GB002342.

243 Kalnay, E., et al. (1996), The NCEP/NCAR 40-Year Reanalysis Project. *Bulletin of the*

244 *American Meteorological Society*, 77 (3), 437–471.

245 Keeling, C. D. et al., (2001), Exchanges of atmospheric CO<sub>2</sub> and <sup>13</sup>CO<sub>2</sub> with the terrestrial

246 biosphere and oceans from 1978 to 2000. I. Global aspects, SIO Reference Series, No. 01-

247 06, Scripps Institution of Oceanography, San Diego, 88 pp.

248 Ledwell, J. R., D.J. McGillcuddy Jr., L.A. Anderson (2008), Nutrient flux into an intense deep

249 chlorophyll layer in a mode-water eddy. *Deep-Sea Res. II*, 55, 1139-1160.

250 Marshall, J.C., C. Hill, L. Perelman, and A. Adcroft (1997), Hydrostatic, quasi-hydrostatic and

251 non-hydrostatic ocean modeling, *J. Geophys. Res.*, 102, 5733–5752.

252 McNeil, C. and E. D'Asaro (2007), Parameterization of air-sea gas fluxes at extreme wind

253 speeds. , *J. Mar. Syst.*, 66, 110-121.

254 Perrie, W., W. Zhang, X. Ren, Z. Long and J. Hare (2004), The role of midlatitude storms on air-  
255 sea exchange of CO<sub>2</sub>. *Geophys. Res. Lett.* 31, L09306, doi:10.1029/2003GL019212.

256 Takahashi, T., et al. (2002), Global sea-air CO<sub>2</sub> flux based on climatological surface ocean  
257 pCO<sub>2</sub>, and seasonal biological and temperature effects, *Deep-Sea Res. II*, 49, 1601-1622.

258 Wanninkhof, R. (1992), Relationship between wind speed and gas exchange over the ocean, *J.*  
259 *Geophys Research*, 97, C5, 7373-7382.

260 Wanninkhof, R., A. Olsen, and J. Triñanes (2007), Air-sea CO<sub>2</sub> fluxes in the Caribbean Sea from  
261 2002-2004, *J. Mar. Syst.*, 66, 272-284.

262 Ullman, D., G.A. McKinley, V. Bennington and S. Dutkiewicz, (2008), North Atlantic carbon  
263 cycle response to climate variability, *submitted to Global Biogeochem. Cycles.*

264

264 Figure Captions:

265 Figure 1. Modeled and observed variables at Bermuda (31N, 64W) from 15 May 1995 through  
266 15 December 1995. SST and pCO<sub>2</sub> data from B07 and MLDs calculated from CTD profiles  
267 (downloaded from <http://bats.bios.edu/>) indicated with blue asterisks. Hurricanes Felix, Luis, and  
268 Marilyn are denoted by the gray bars, beginning one day before closest approach to Bermuda and  
269 ending one day after: (a) NCEP reanalysis 10-meter wind speeds, (b) MLD, (c) SST, (d) Total  
270 pCO<sub>2</sub> (thin solid bold black), pCO<sub>2</sub>-T (thin red), pCO<sub>2</sub>-nontT (dashed red), and pCO<sub>2</sub><sup>atm</sup> (thin  
271 black), and (e) air-sea CO<sub>2</sub> flux, positive to the atmosphere.

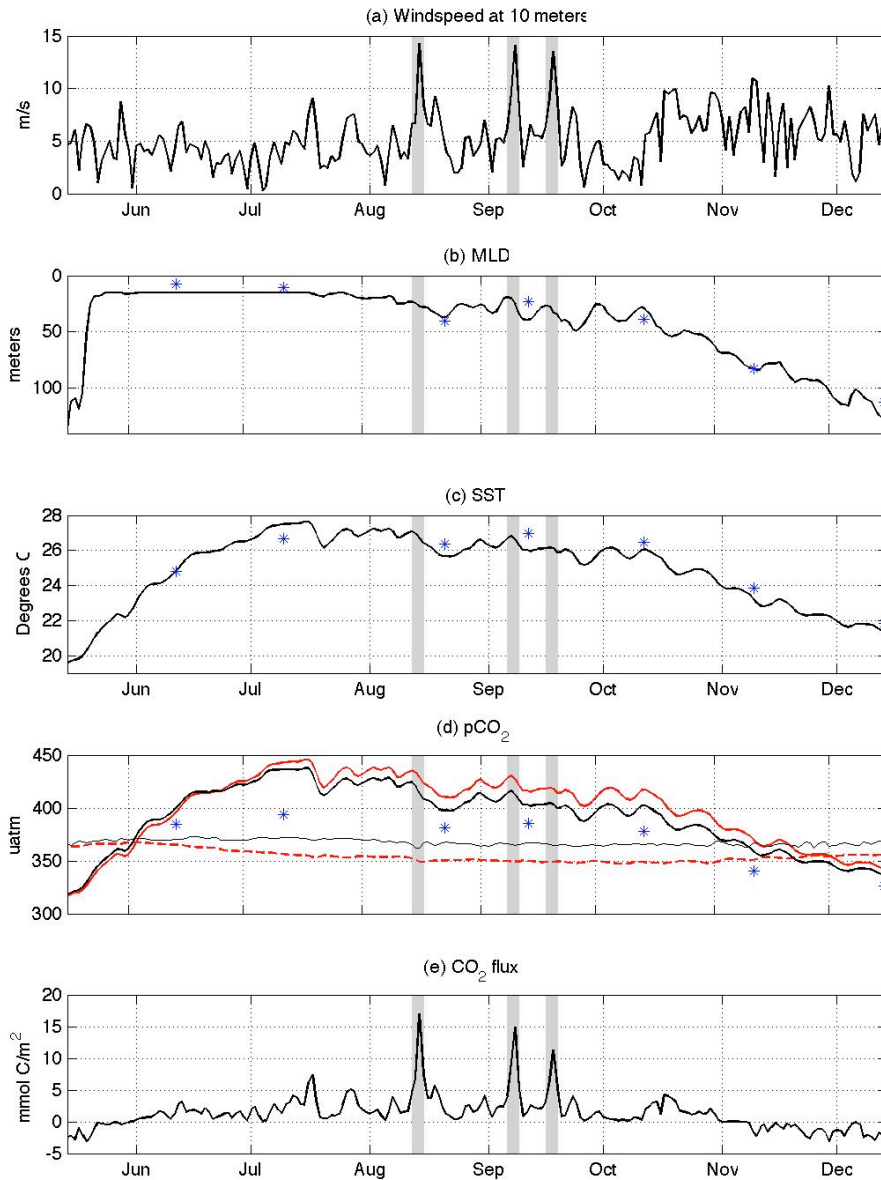
272

273 Figure 2. Profiles of (top) phosphate and (bottom) DIC concentrations from the model. Dashed  
274 lines are from a day before the hurricanes passed Bermuda, solid lines from a day after. On the  
275 phosphate profiles, the depth of the modeled MLDs the day before (triangle) and the day after  
276 (square) the hurricane are also indicated. Asterisks indicate the observed monthly nutrient  
277 profile.

278

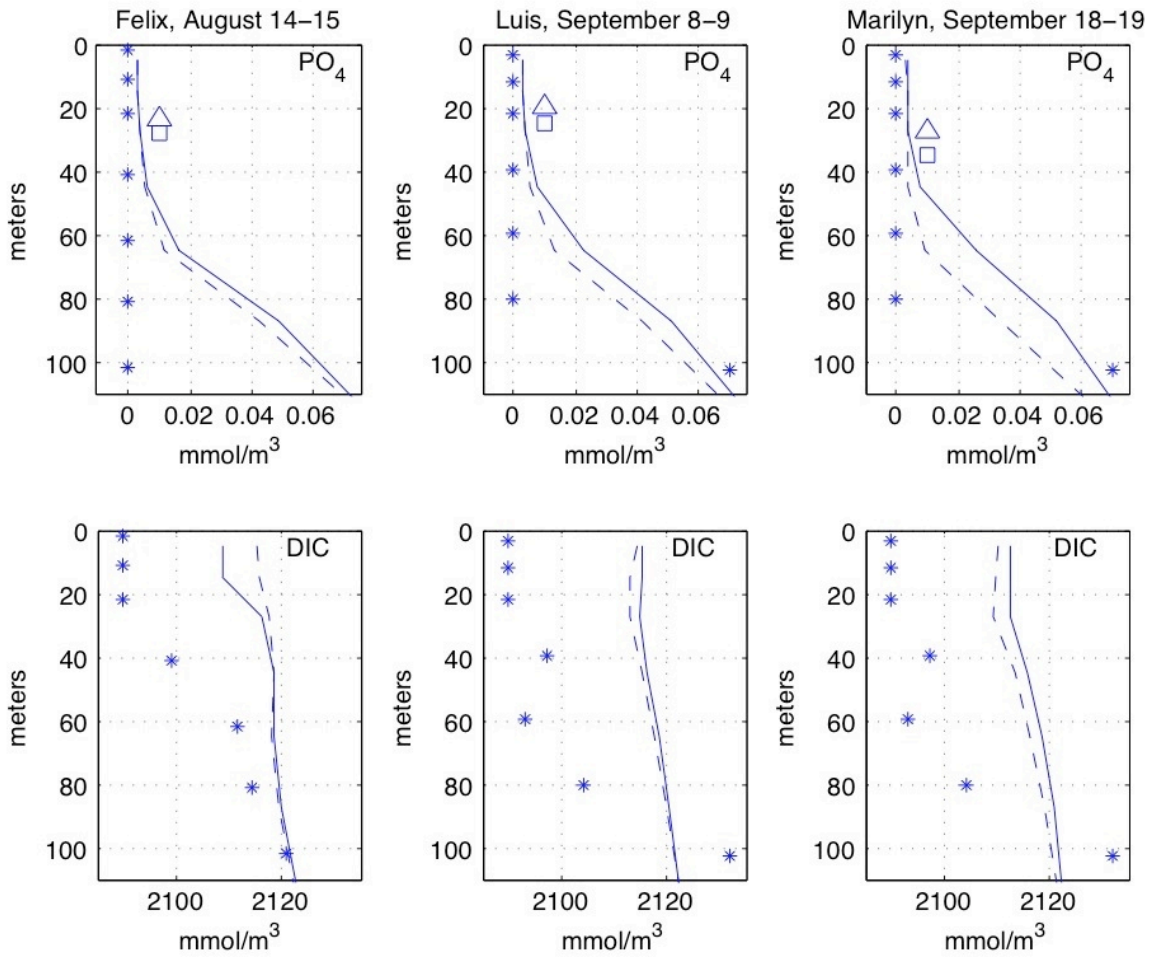
279 Figure 3. Scatter plots of year, hurricane and tropical storm (TS) frequency and modeled CO<sub>2</sub>  
280 fluxes, as well as best-fit lines when a linear trend is able to reasonably define the relationship.  
281 (a) Hurricane and tropical storm frequency (NHC data) versus year, (b) annual flux (25-40N,  
282 including the Gulf of Mexico and Caribbean) per year, (c) sum of the daily fluxes along  
283 hurricane and tropical storm tracks (NHC data) versus number of hurricanes and tropical storms,  
284 (d) annual flux versus hurricane and tropical storm frequency.

285



287

288 Figure 1. Modeled and observed variables at Bermuda (31N, 64W) from 15 May 1995 through  
 289 15 December 1995. SST and pCO<sub>2</sub> data from B07, and MLDs calculated from CTD profiles  
 290 (downloaded from <http://bats.bios.edu/>) indicated with blue asterisks. Hurricanes Felix, Luis, and  
 291 Marilyn are denoted by the gray bars, beginning one day before closest approach to Bermuda and  
 292 ending one day after: (a) NCEP reanalysis 10-meter wind speeds, (b) MLD, (c) SST, (d) Total  
 293 pCO<sub>2</sub> (thin solid bold black), pCO<sub>2</sub>-T (thin red), pCO<sub>2</sub>-nontT (dashed red), and pCO<sub>2</sub><sup>atm</sup> (thin  
 294 black), and (e) air-sea CO<sub>2</sub> flux, positive to the atmosphere.

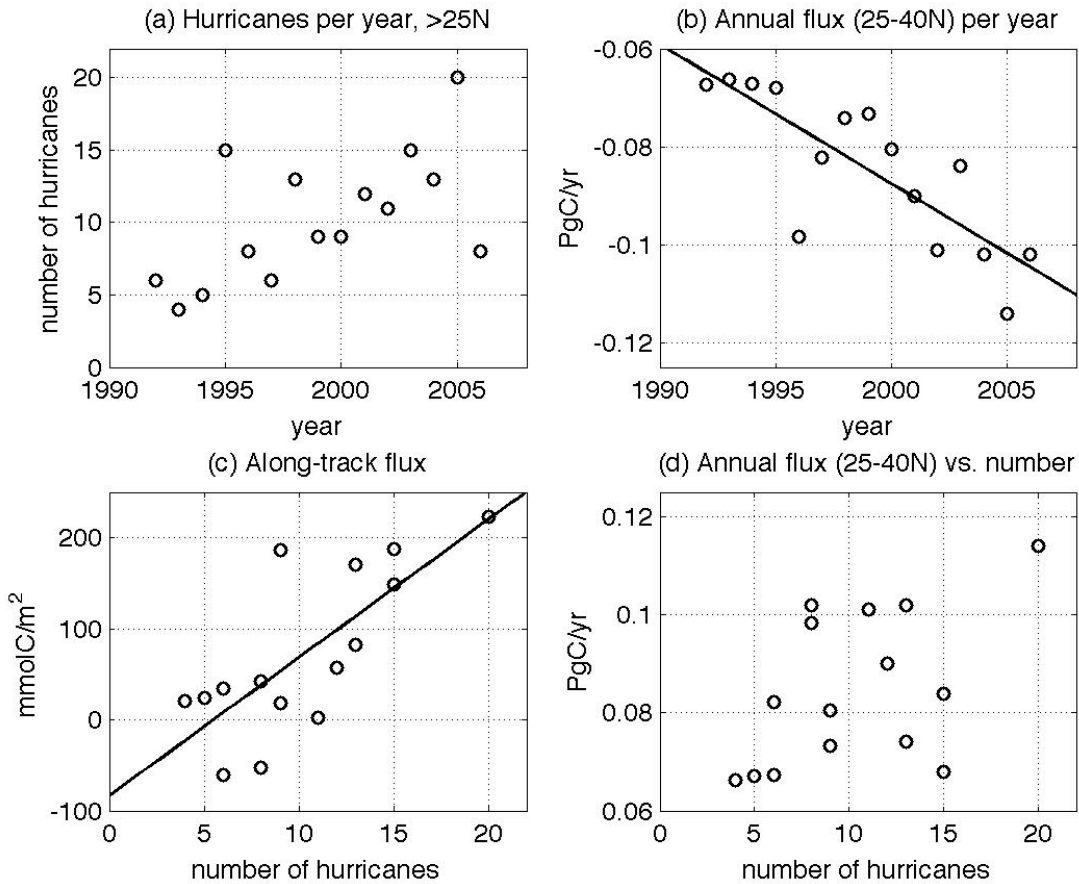


295

296 Figure 2. Profiles of (top) phosphate and (bottom) DIC concentrations from the model. Dashed  
 297 lines are from a day before the hurricanes passed Bermuda, solid lines from a day after. On the  
 298 phosphate profiles, the depth of the modeled MLDs the day before (triangle) and the day after  
 299 (square) the hurricane are also indicated. Asterisks indicate the observed monthly nutrient  
 300 profile.

301

302



302  
 303 Figure 3. Scatter plots of year, hurricane and tropical storm (TS) frequency and modeled CO<sub>2</sub>  
 304 fluxes, as well as best-fit lines when a linear trend is able to reasonably define the relationship.  
 305 (a) Hurricane and tropical storm frequency (NHC data) versus year, (b) annual flux (25-40N,  
 306 including the Gulf of Mexico and Caribbean) per year, (c) sum of the daily fluxes along  
 307 hurricane and tropical storm tracks (NHC data) versus number of hurricanes and tropical storms,  
 308 annual flux versus hurricane and tropical storm frequency. Positive fluxes are to the  
 309 atmosphere.  
 310